

# Seasonality of $^7\text{Be}$ concentrations in Europe and influence of tropopause height

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## ABSTRACT

This study aims at analysing the latitudinal variability of both the yearly and seasonal pattern of  $^7\text{Be}$  surface activity concentrations, at addressing the impact of tropopause height (TPH) on  $^7\text{Be}$  distribution and at evaluating the time lag between TPH and  $^7\text{Be}$  at European level. With this aim, weekly  $^7\text{Be}$  and daily TPH data at 17 sampling stations during 10 yr (2001–2010) are analysed.  $^7\text{Be}$  shows a clear increasing tendency in the period and generally tends to increase with decreasing latitude. The seasonal pattern generally shows maxima during the warm period and minima during the cold one. The seasonal variogram analysis points out a good spatial correlation for TPH data while a weaker one is observed for  $^7\text{Be}$ , having TPH a larger influence on  $^7\text{Be}$  during summer. The influence of TPH on  $^7\text{Be}$  exhibits a large spatial variability, with a clear gap between south and north in the area of the polar front jet. The results identify the presence of two main groups, in particular separating between stations located in northern Europe (50 °N and higher) and stations in southern Europe (south of 50 °N). A similar behaviour for stations located in the same geographical area is also observed when looking at the day of maximum impact of TPH on  $^7\text{Be}$  concentrations. The results suggest that  $^7\text{Be}$  concentrations respond in different time ranges to changes in the TPH, observing seasonal differences in each group. These results represent the first European approach to the understanding of the TPH impact on  $^7\text{Be}$  concentrations at surface levels.

*Keywords:*  $^7\text{Be}$ , radiotracers, tropopause height, transport time, seasonality, Europe

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## Key Points

- Variability of  $^7\text{Be}$  activity concentration in relation to tropopause height is investigated.
- Southern and northern areas with different spatial variability are identified.
- Different response times of  $^7\text{Be}$  concentration to changes in tropopause height are found.

## 1. Introduction

Beryllium-7 ( $^7\text{Be}$ ) cosmogenic isotope is mainly (75 %) produced in the stratosphere, while the remainder 25 % is produced in the upper troposphere (Usoskin and Kovaltsov, 2008). The residence time of  $^7\text{Be}$  in the stratosphere is typically of the order of a couple of years (e.g. McHargue and Damon, 1991). Once produced,  $^7\text{Be}$  is rapidly adsorbed onto fine aerosol particles in the stratosphere and further transported to the troposphere. Once in this layer, several factors affect the distribution of  $^7\text{Be}$  in surface air, introducing complexity in the evaluation of the  $^7\text{Be}$  activity in

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surface air (Baskaran, 2011), being estimated mean residence times in the troposphere from 18 to 23 d (e.g. Koch et al., 1996). However, as a first step in this complex analysis, the so-called stratosphere-to-troposphere transport (STT) is one of the key factors to understand  $^7\text{Be}$  activity concentrations at surface levels.

The life cycle of STT including the decay, dispersion and eventually mixing of the air masses has been studied (e.g. Cooper et al., 2004). In this sense, previous studies using, among others, also  $^7\text{Be}$  as a tracer of stratospheric air in the troposphere (Zanis et al., 1999; Cristofanelli et al., 2003, 2006) have demonstrated that air parcels originating in the stratosphere may reach the surface levels in two ways: (1) directly (less than about 2 d), through a vigorous vertical transport, and with largely conserved chemical–physical properties after intrusion into the troposphere, and (2) indirectly (about 4–5 d), through a multistep transport process, and a greater loss of stratospheric properties along the path. The latter one, the indirect transport, may affect air composition 1000 km away from the intrusion area.

As the occurrence of STT can be connected with intense tropopause deformations (e.g. Reid and Vaughan, 2004; Cristofanelli et al., 2006), the tropopause height (TPH) evolution can be used to identify the location and the characteristics of STT. The TPH is characterised by both a large zonal variation, reaching values of 15–18 km in the tropics and no more than 6–8 km at the poles (Liu et al., 2014), as well as a temporal variability, presenting a seasonal and a daily variation linked with the presence of weather systems. Within the temporal and space variability of the TPH, many studies have focused on the occurrence and characteristics of tropopause folds related to cyclogenesis at midlatitudes (e.g. Browning, 1997) and resulting in deep intrusions of stratospheric air into the troposphere (e.g. Reid and Vaughan, 2004).

Some studies have focused on the influence of STT on  $^7\text{Be}$  activity concentrations at surface level during tropopause folding events. For example, Gerasopoulos et al. (2006) analysed a complex case study of STT in Greece in which  $^7\text{Be}$  concentrations exceeded  $9\text{ mBq m}^{-3}$ ; Stohl et al. (2000) analysed  $^7\text{Be}$  concentrations (and ozone and humidity) to study frequency of stratospheric intrusions in the Alps and northern Apennines; and Zanis et al. (2003) investigated a predicted deep STT event encountered during the STACCATO project period on 20–21 June 2001. In addition, Ioannidou et al. (2014) provided an estimation of the time lag between the TPH and surface  $^7\text{Be}$  concentrations.

These studies represent sporadic previous evidences for certain locations in Europe, but a comprehensive analysis of the impact of TPH on  $^7\text{Be}$  distribution across Europe has not yet been undertaken. This kind of analysis is crucial to better understand the regional distribution of  $^7\text{Be}$ ; moreover, the results can be useful in establishing the

relationship between  $^7\text{Be}$  and meteorological factors (such as dry and wet deposition, synoptic and mesoscale circulations) which affect its concentrations at surface levels.

Therefore, taking the TPH evolution as an indicator of STT, the present study aims at: (1) offering a better understanding of the STT impact on the  $^7\text{Be}$  activity concentrations at surface levels, and (2) providing an estimation of the time lag between the elevation of TPH and surface  $^7\text{Be}$  concentrations at European level.

## 2. Material and methods

### 2.1. $^7\text{Be}$ database: monitoring sites

The current study used  $^7\text{Be}$  data collected between 2001 and 2010 at the sampling locations indicated in Table 1 and shown in Fig. 1. These data were stored by the online Radioactivity Environmental Monitoring (REM) Database, supported by the Radioactivity Environmental Monitoring group of the Institute for Transuranium Elements (ITU) of the Joint Research Centre (JRC). Nowadays, REMdb contains a unique collection of environmental radioactivity measurements from 1984 onwards to analyse and compare environmental radioactivity data in Europe ([www.rem.jrc.ec.europa.eu/RemWeb](http://www.rem.jrc.ec.europa.eu/RemWeb)). The  $^7\text{Be}$  database stored in the REMdb is public until 2006, while the access to the data corresponding to the 2007–2011 period can be granted only after explicit request. The groups providing the data have already published several papers on their radioactivity measurements [e.g. Dueñas et al. (2011) and Piñero Garcia et al. (2012) in Spain; Leppänen and Paatero (2013) in Finland; Ioannidou et al. (2005) in Greece; and Carvalho et al. (2013) in Portugal].

The period of 2001–2010 was selected to provide the highest number of operative stations with available measurements. Individual radionuclide analyses extracted from REMdb were mainly performed on a weekly basis, with a varying time resolution from 6 to 8 d. Table 1 also shows the number of measurements used in this work.  $^7\text{Be}$  was measured on airborne particulate matter filters sampled through high-volume particulate samplers. Further information on the procedure, as well as the list of National Competent Authorities, is available in Hernández-Ceballos et al. (2015).

### 2.2. TPH calculation

Daily TPH values were calculated at each station following the methodology of Ioannidou et al. (2014). This method is based on the use of values of geopotential height and pressure on the closest isobaric levels (above and below the tropopause) and the air temperature and pressure at the tropopause level. Data used for the calculation were downloaded from the Earth System Research Laboratory (ESRL).

Table 1. Station name, coordinates (latitude, longitude, altitude above sea level), number of operability sampling periods and temporal resolution.

Stations	LAT (°N)	LON (°E)	Altitude (m a.s.l)	Operability (samplings during 2001–2010)	Temporal resolution
Ivalo	68.64	27.57	130	418	Weekly (between 6 and 8 d)
Umea	63.85	20.34	45	478	
Kista	59.4	17.93	16	488	
Harku	59.39	24.58	36	400	
Risoe	55.69	12.10	9	448	
Bilthoven	52.11	5.18	4	501	
Offenbach	50.1	8.77	107	445	
Praha	50.09	14.42	202	446	
Luxembourg	49.63	6.13	280	434	
Vienna	48.22	16.35	193	506	
Freiburg	48.20	7.87	411	466	
Milano	45.47	9.18	125	349	
Bilbao	43.17	-2.94	380	468	
Barcelona	41.38	2.12	52	472	
Madrid	40.45	-3.69	715	471	
Sevilla	37.39	-6.01	8	475	
Laguna	28.46	-16.29	358	200	

### 2.3. Statistical methods

Geostatistics offers a way to describe the spatial continuity of natural phenomena and provides adaptations of classical regression techniques to take advantage of this continuity

(Isaaks and Srivastava, 1989). Therefore, geostatistical tools were used to study the spatial structure of  $^7\text{Be}$  and TPH data. To understand the spatial variability of  $^7\text{Be}$  and TPH in Europe and in particular, to infer how much of the  $^7\text{Be}$  variability can be explained by the location (longitude,

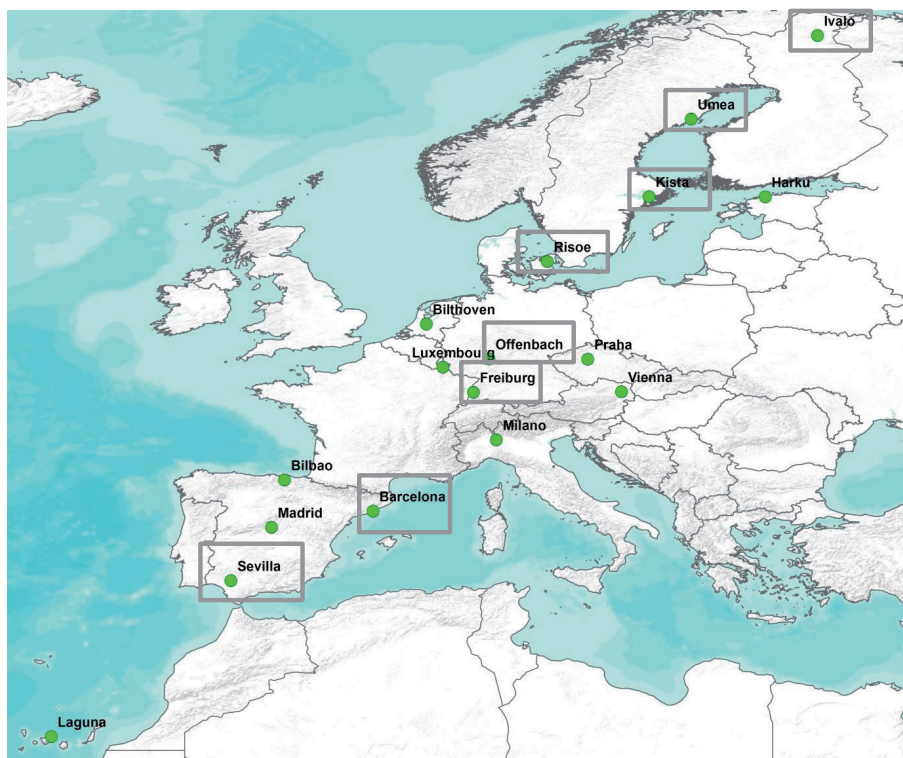


Fig. 1. Map of  $^7\text{Be}$  sampling sites. The stations given in squares are selected for latitudinal analysis (Section 3.2).

latitude) or by other factors, the spatial correlation between data was measured by means of the variogram. Variograms were calculated using the software Surfer11 (SURFER 2012), considering the average  $^7\text{Be}$  concentrations and TPH over the period 2001–2010 and looking separately into winter (Dec–Jan–Feb) and summer (Jun–Jul–Aug) seasons.

The distribution of data in each station was tested. In the majority of the stations, the normality test was verified. The Pearson's linear correlation coefficient (significance level 0.05) was calculated for each selected sampling site to determine the degree of association between TPH and  $^7\text{Be}$  activity concentrations. In all the statistical analyses, spring and autumn were not considered since they are transition seasons.

### 3. Results and discussion

#### 3.1. Variogram analysis of TPH and $^7\text{Be}$

Variograms for  $^7\text{Be}$  and TPH are presented in Fig. 2a–f. The x-axis represents the distance between the pairs of samples while the value of the experimental variogram is plotted along the y-axis. In the present study, the  $^7\text{Be}$  variograms did not present a regular trend; that is, two observations closer together are not more similar than two observations farther apart – it means there is not a strong spatial correlation neither during the total period (Fig. 2a) nor during the winter–summer seasons (Fig. 2c and e). Variograms for TPH showed instead a good spatial correlation only at short distances considering the total database (Fig. 2b), while no spatial correlation was observed for TPH during winter (Fig. 2d). The summer variogram for TPH (Fig. 2f) gave instead a high spatial correlation, which was even higher than that observed for the total period since it was maintained also at larger distances.

This analysis suggested that, as expected, TPH exerts only a minor influence on  $^7\text{Be}$  concentrations, and other factors, such as rainfall or wind dynamics, may strongly affect  $^7\text{Be}$  activity concentrations at surface levels especially during the winter season. In fact, the exchange between stratosphere and troposphere is only about the 20 % of the total amount of  $^7\text{Be}$  measured at surface level (Dutkiewicz and Husain, 1985). The analysis also highlighted the importance of local/regional/synoptic surface conditions on  $^7\text{Be}$  activity concentrations. It is worth noting here that all the stations in the REMdb are located at low altitude and within the planetary boundary layer (and for this reason local conditions might play a key role). The important role of local/mesoscale meteorological conditions on  $^7\text{Be}$  concentrations probably causes the absence of a spatial correlation for this variable.

At this point and in the light of the spatial correlation of TPH observed, we decided to select a set of stations to analyse the potential impact of TPH latitudinal variability on  $^7\text{Be}$  activity concentrations. Hence, a set of eight sampling locations along an imaginary line from low (Sevilla, as Laguna presented less variability of data) to high (Ivalo) latitudes were selected (the selected stations are marked with squares in Fig. 1).

#### 3.2. Yearly and monthly variability of $^7\text{Be}$

Annual average  $^7\text{Be}$  concentrations were calculated for the selected set of networks. As shown in Fig. 1 and as discussed in the previous section, these stations were selected to address the latitudinal variability of  $^7\text{Be}$  in Europe. Figure 3 shows the spatial (from low to high latitudes) and temporal (from 2001 to 2010) variation of the annual arithmetic means of the weekly measurements.

As a first remark, a clear interannual increasing tendency could be noticed in all the stations; moreover, a generally decreasing mean  $^7\text{Be}$  value with increasing latitude was also noticed (we refer to Hernandez-Ceballos et al., 2015, which gives a more general overview of  $^7\text{Be}$  activity concentrations in Europe). This interannual pattern confirmed the impact of the 11-yr solar modulation on  $^7\text{Be}$  air concentrations (Usoskin and Kovaltsov, 2008; Leppänen et al., 2010). The first tendency was a result of the  $^7\text{Be}$  production mechanism, which varies with solar modulation of galactic cosmic rays invading the heliosphere (Masarik and Beer, 1999), and in turn, is controlled by solar activity. The evolution of the sunspot number ([www.sidc.oma.be/silso/datafiles](http://www.sidc.oma.be/silso/datafiles)) showed that our reference period (2001–2010) encompasses a part of the 23rd (1996 May–2008 January) and the beginning of the 24th (2008 January – present) solar cycle. An analysis of the annual mean  $^7\text{Be}$  concentrations with sunspot number clearly showed the opposite behaviour (Kikuchi et al., 2009; Steinmann et al., 2013) (not shown).

In the second case, the latitudinal influence on  $^7\text{Be}$  production (Kulan, 2006) caused a general decrease in the yearly values with an increase in latitude of the considered stations, resulting in higher values in the southern than in northern sampling stations. The minimum concentration was found in Ivalo (1.9 mBq/m<sup>3</sup>), while the maximum one was found in Sevilla (4.1 mBq/m<sup>3</sup>).

The seasonal variability of  $^7\text{Be}$  concentration at surface levels is displayed in Fig. 4, showing the presence of a strong seasonal pattern in all the stations: generally speaking, increased concentrations during the warm period from May to August, with a maximum during the months of June–July, and lower concentrations in winter and autumn, were observed. This pattern was consistent with previous findings for  $^7\text{Be}$  concentrations at surface levels (e.g.

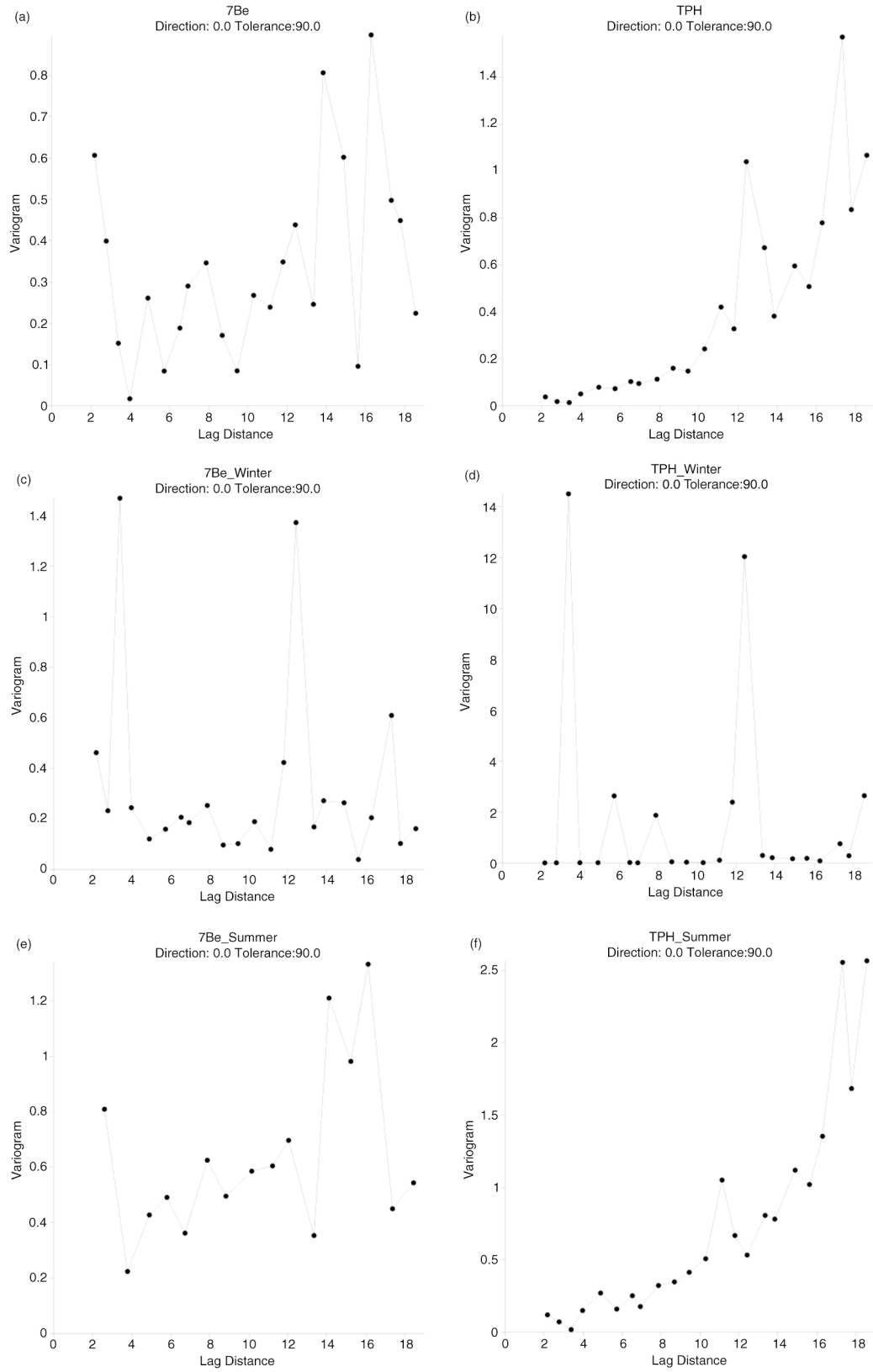


Fig. 2. Variograms for  $^7\text{Be}$  and TPH: all available data (a and b), and separately winter (c and d) and summer (e and f) seasons.

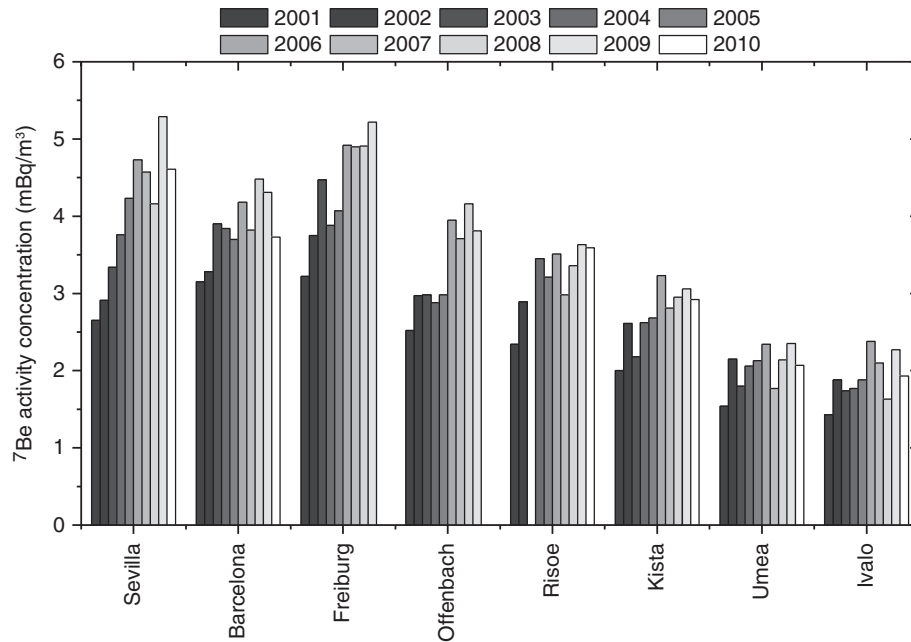


Fig. 3. Interannual variability of <sup>7</sup>Be activity concentrations.

Lozano et al., 2011; Tositti et al., 2014). This seasonal variation could be explained by a combination of several factors, among which is the increased vertical transport of <sup>7</sup>Be from the upper troposphere to the middle and lower troposphere in the summer months (e.g. Feely et al., 1982, 1989). In addition, similarly to the findings for the annual means, the influence of latitude on <sup>7</sup>Be activity concentrations was also noted in this case.

### 3.3. Variability of TPH in EU: Correlation with monthly <sup>7</sup>Be

Tropopause is recognised as a key feature of the atmospheric structure at all latitudes (Seidel and Randel, 2006). Ioannidou et al. (2014) demonstrated a strong positive correlation between the seasonal changes of <sup>7</sup>Be concentration in surface air and TPH and temperature ( $R = 0.97$  and  $R = 0.94$ , respectively) in Thessaloniki, Greece.

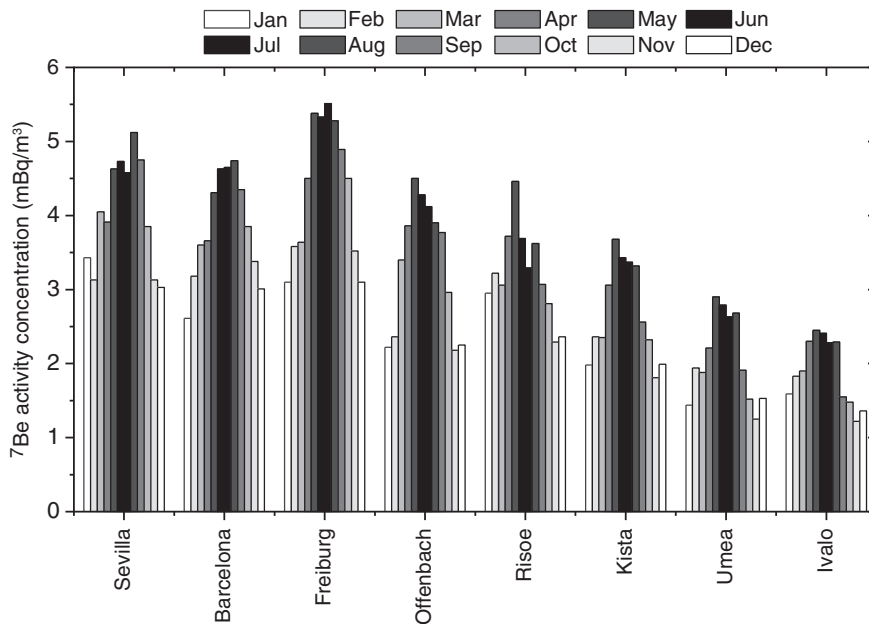


Fig. 4. Mean monthly variability of <sup>7</sup>Be activity concentrations (2001–2010).

Figure 5 shows the correlation coefficient obtained between the <sup>7</sup>Be and TPH monthly means over the 2001–2010 period (12 monthly values for each variable and for each sampling station). A comparison of the coefficients indicated a large spatial variability in the impact of TPH on <sup>7</sup>Be activity concentrations. According to the correlation coefficient value, two clear groups arose (from right to left in Fig. 5). The first group comprehended stations with a low (less or equal to 0.4) and even negative (Risoe and Ivalo) value of the correlation coefficient. These measurement sites were located in the northern-central Europe (50 °N and higher). The second group was formed by stations with a higher value of the correlation coefficient (up to 0.6) which were located in the southern-central Europe (south of 50 °N).

Within the first group of the measurement sites (north of 50 °N), Risoe and Ivalo, which are located in the vicinity of the Atlantic Ocean, exhibited a different value of the correlation coefficient (Fig. 5); in fact, the correlation coefficient for these two stations was negative, unlike the coefficients for all other sites. This observation might result from the fact that these sites are strongly affected by maritime air masses transported by the westerlies, which have low concentration of aerosols (Sorribas et al., 2011), and therefore, low concentration of <sup>7</sup>Be attached to them. A similar observation was reported by Gerasopoulos et al. (2001). It is worth noting here that these exceptions along the Atlantic Ocean coast imply a limitation of our analysis, in which ‘a horizontally stationary’ troposphere is assumed, which will be further discussed below.

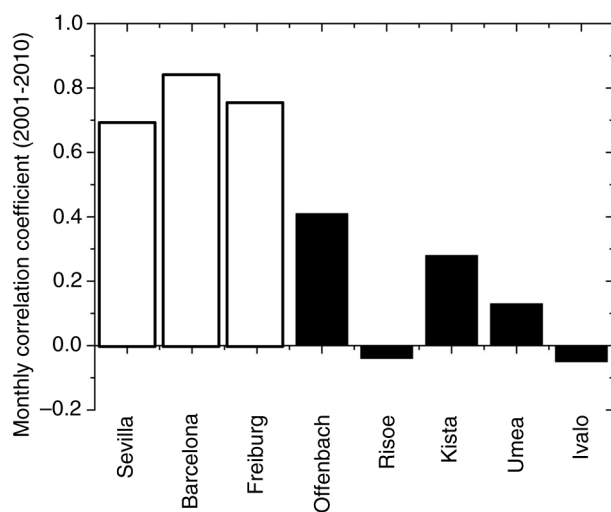


Fig. 5. Correlation coefficient between mean monthly <sup>7</sup>Be and TPH (2001–2010). White and black columns denote significant and non-significant correlations, respectively.

These results displayed the existence of a latitudinal variability in the impact of TPH on <sup>7</sup>Be activity concentrations. As reported above in the Introduction, TPH shows a latitudinal and a seasonal dependence (Liu et al., 2014), and it is higher over the tropics (~16 km) than at the poles (~8 km). However, the TPH decrease from low to high latitudes is not gradual – there is a rapid drop in TPH in the area of the polar front jet (50–60 °N). This behaviour is mostly due to the polar jet stream, a relatively narrow band of strong wind in the upper levels of the atmosphere. In fact, TPH suddenly decreases across the jet stream especially in conditions of strong jet streams associated with a front at lower levels. For this reason, the area of the polar front jet represents a gap in the gradual latitudinal decrease of TPH towards the poles, and this behaviour may partially explain the spatial variability between TPH and <sup>7</sup>Be activity concentrations.

To confirm this relationship, Fig. 6 shows the monthly evolution of TPH and <sup>7</sup>Be activity concentrations in the eight sampling locations over the 2001–2010 period. The scatter plot diagram of <sup>7</sup>Be vs. TPH for each station can be consulted in supporting information (Supplementary Fig. 1), which reports a coefficient of determination ( $R^2$ ) less than 0.4, reaching the highest values in the southern-central stations and the lowest ones in the northern ones. The latitudinal decrease was observed in both variables, and specifically for TPH from a minimum of six kilometres at Ivalo to a maximum of about nine kilometres in Seville. In addition to spatial differences, the monthly cycle was also pronounced, with a maximum in the warm months and a minimum during the cold months. This variability was mainly associated with the cycle of solar irradiation as well as with changes in meteorological conditions (e.g. Lee et al., 2007; Usoskin and Kovaltsov, 2008).

### 3.4. Spatial and temporal variability of time lag

The time lag between TPH and levels of <sup>7</sup>Be surface concentrations represents the time required for the <sup>7</sup>Be concentrations levels in the near surface air to respond to the change of TPH (Ioannidou et al., 2014). Given the spatial and temporal variability of the TPH, this time lag can register large spatial, from one station to other, and temporal, from hours to days, differences.

A correlation between weekly <sup>7</sup>Be values and the corresponding seven daily TPH values was then calculated to describe and analyse the behaviour of this relationship across Europe. For each location, Pearson’s linear correlation coefficients between weekly <sup>7</sup>Be and daily TPH were calculated with a time lag of  $n$  days. The number of days  $n$  was allowed to vary from zero (corresponding to correlating the weekly <sup>7</sup>Be activity concentration with the TPH on the sampling end day) to seven (corresponding to

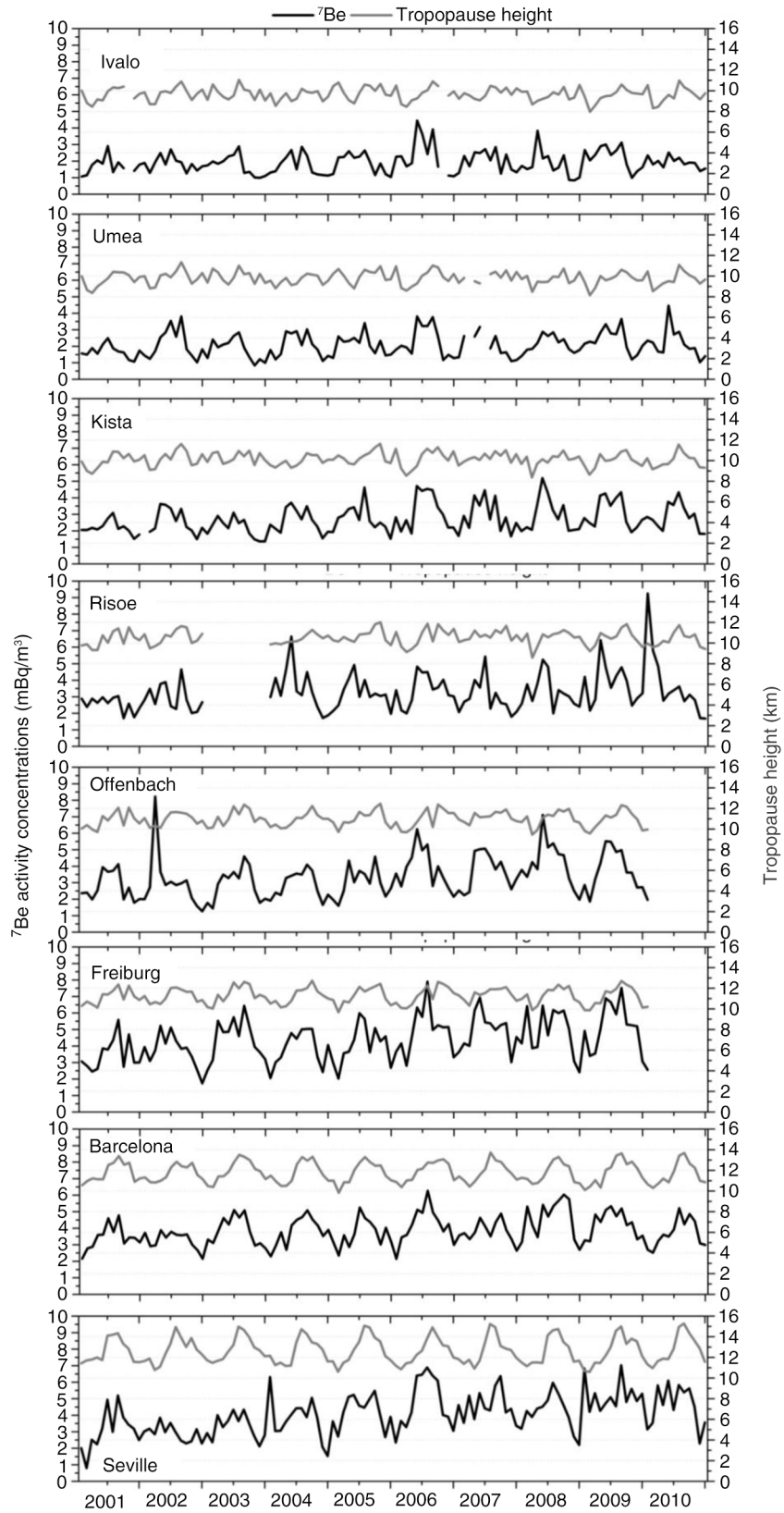


Fig. 6. Latitudinal variation of the monthly evolution of  $^7\text{Be}$  and TPH (2001–2010).



correlating the weekly <sup>7</sup>Be activity concentration with the TPH on the sampling start day).

The selection of this maximum value of *n* (time lag period) was based on previous works (e.g. Cristofanelli et al., 2006), indicating that air parcels originating in the stratosphere may reach the surface within 5 d. In addition, as our data were weekly <sup>7</sup>Be values, the use of weekly or average TPH values to calculate the correlation would generate a smooth component in the analysis, affecting the correlation values and hence, the interpretation of the results.

This methodology offered information on: (1) the temporal variability of the TPH impact on weekly <sup>7</sup>Be activity concentration (maximum correlation coefficient) in each station, and (2) the spatial variability of the time lag.

The results for the eight selected stations are presented in Fig. 7a, showing the pattern of the cross-correlation analysis of weekly <sup>7</sup>Be values and daily TPH. The maximum correlation values found in the analysis ranged between 0.2 and 0.45, which is in agreement with the 0.44 value reported by Ioannidou et al. (2014). The time lag corresponding to the maximum coefficient value mainly

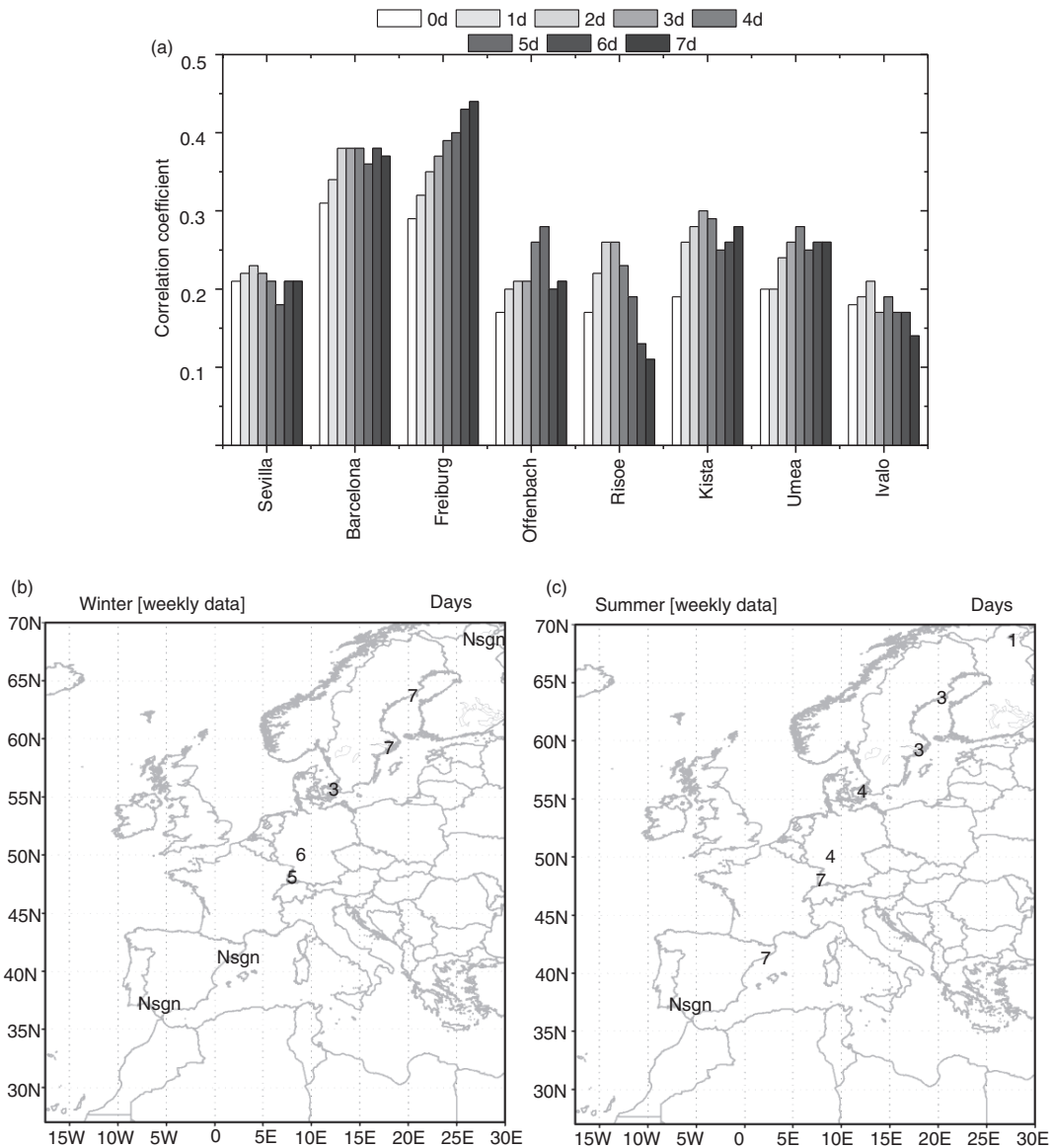


Fig. 7. (a) Correlogram (a time lag from 0 to 7 d) between <sup>7</sup>Be surface concentration and tropopause height, and maps showing the day on which the maximum correlation coefficient between TPH and <sup>7</sup>Be is observed at the selected stations during (b) winter and (c) summer. 'Nsgn' stands for a correlation coefficient not statistically significant.

ranged between 3 and 5 d. An exception was Freiburg where the correlation coefficient registered the maximum (0.45) on the 7th day. This behaviour highlighted the large spatial variation of the time delay between the TPH and  $^7\text{Be}$  concentrations in surface air. In the southern stations (Sevilla and Barcelona), the correlation coefficients had similar values over the days and after reaching the maximum value (on the third day), a plateau was established. Conversely, at higher latitude stations, more marked daily differences between the correlation values prior to and after reaching the maximum value were observed. In this line, it is worth pointing out the sharp decrease of the correlation value observed in Risoe, which might be associated with the persistence time in the state of the atmosphere. Finally, no particular tendency (neither to increase nor to decrease) for the day of the maximum correlation with latitude was detected.

The comparison of the day in which the maximum correlation coefficient was reached in winter and summer (Fig. 7b and c) pointed out additional differences between the measurement sites. While in the northern stations this day decreased from winter to summer (from up to 7 d down to less than or equal to 3 d), in the central area generally higher spatial variability among the stations, but with reduced differences in magnitude, were observed. For example, from winter to summer it increased in Risoe (from three to four) and Freiburg (from five to seven), but decreased in Offenbach (from six to four). These patterns might be connected with the seasonal spatial distribution of STT in the northern Hemisphere (Sprenger and Wernli, 2003) and the synoptic configuration associated with it.

As for the southern stations, the correlation was not statistically significant in winter, which indicated the reduced impact of TPH on the  $^7\text{Be}$  concentrations. The result obtained in summer for Barcelona (time lag of 7 d) might be linked with the elevated occurrence of regional recirculations at lower levels over this area (Jorba et al., 2004), which might limit the direct impact of STT on  $^7\text{Be}$  activity concentrations.

These annual and winter–summer patterns for the time lag in the selected stations were also observed when considering all the stations (Supplementary Fig. 2), where it was possible to identify a similar behaviour in the same geographical area.

The comparison of the winter and summer correlation coefficients between  $^7\text{Be}$  surface concentrations and TPH calculated in each station for time lag ranging from 0 to 7 d is shown in Fig. 8a and b. The results showed that the number of significant correlations increased from winter to summer; moreover, the values of the correlation coefficient generally tend to increase from winter to summer, although with some exceptions (Freiburg with 3 d of time lag decreasing from winter to summer from 0.26 to 0.23, and

Offenbach with 5 d of time lag decreasing from winter to summer from 0.30 to 0.24). These results confirmed that the influence of TPH on  $^7\text{Be}$  concentrations was higher during summer than during winter, in agreement with previous studies (e.g. Ioannidou et al., 2014). In contrast to the total set of data, when considering winter and summer seasons separately, the correlation appeared to be generally higher in the central stations, somewhat lower in the northern stations, and practically null in the southern stations.

As already mentioned in the discussion of the negative correlation coefficients obtained for Risoe and Ivalo (Fig. 5), they might be an indication of the impact the horizontal transport within the troposphere exerts on  $^7\text{Be}$  activity concentrations. Horizontal transport can play an important role in the distribution of  $^7\text{Be}$  in surface air (Feely et al., 1989; Gerasopoulos et al., 2001). It is, therefore, important to note here that an indirect exclusion of horizontal transport within the troposphere introduced a limitation in our analysis. Hence, in the future, a more comprehensive study, including both vertical and horizontal transport in the troposphere, would undoubtedly give more insight into the relation between the STT and  $^7\text{Be}$  distribution in surface air.

## 4. Conclusions

Ten years of  $^7\text{Be}$  data in 17 sampling stations along Europe, stored in the REMdb, were analysed together with the TPH in order to characterise: (1) the spatial and temporal distribution of  $^7\text{Be}$  in surface air, (2) the impact of the TPH on  $^7\text{Be}$  surface activity concentration, and (3) the time lag variability between the TPH and  $^7\text{Be}$  concentration.

Regarding the first objective, the  $^7\text{Be}$  activity concentrations showed both spatial and temporal variability. The  $^7\text{Be}$  activity concentration increased as latitude decreased from the north to south (as shown by the variogram analysis). On the other hand, the temporal variability was marked by a clear interannual tendency, governed by the solar cycle, and a noticeable seasonal pattern, with maxima during the warm months and minima during winter and autumn, most likely associated with a combination of meteorological factors.

In the analysis of the relationship between  $^7\text{Be}$  and TPH, the variogram results showed that TPH exerted only a minor influence on the  $^7\text{Be}$  activity concentrations during the winter season, and a larger influence in summer. In general, the results displayed latitudinal variability in the impact of TPH on  $^7\text{Be}$  activity concentrations, also reflecting a gap in the gradual latitudinal decrease of TPH towards the North Pole in the area of the polar front jet. The spatial distribution of the time lag revealed a spatial and temporal

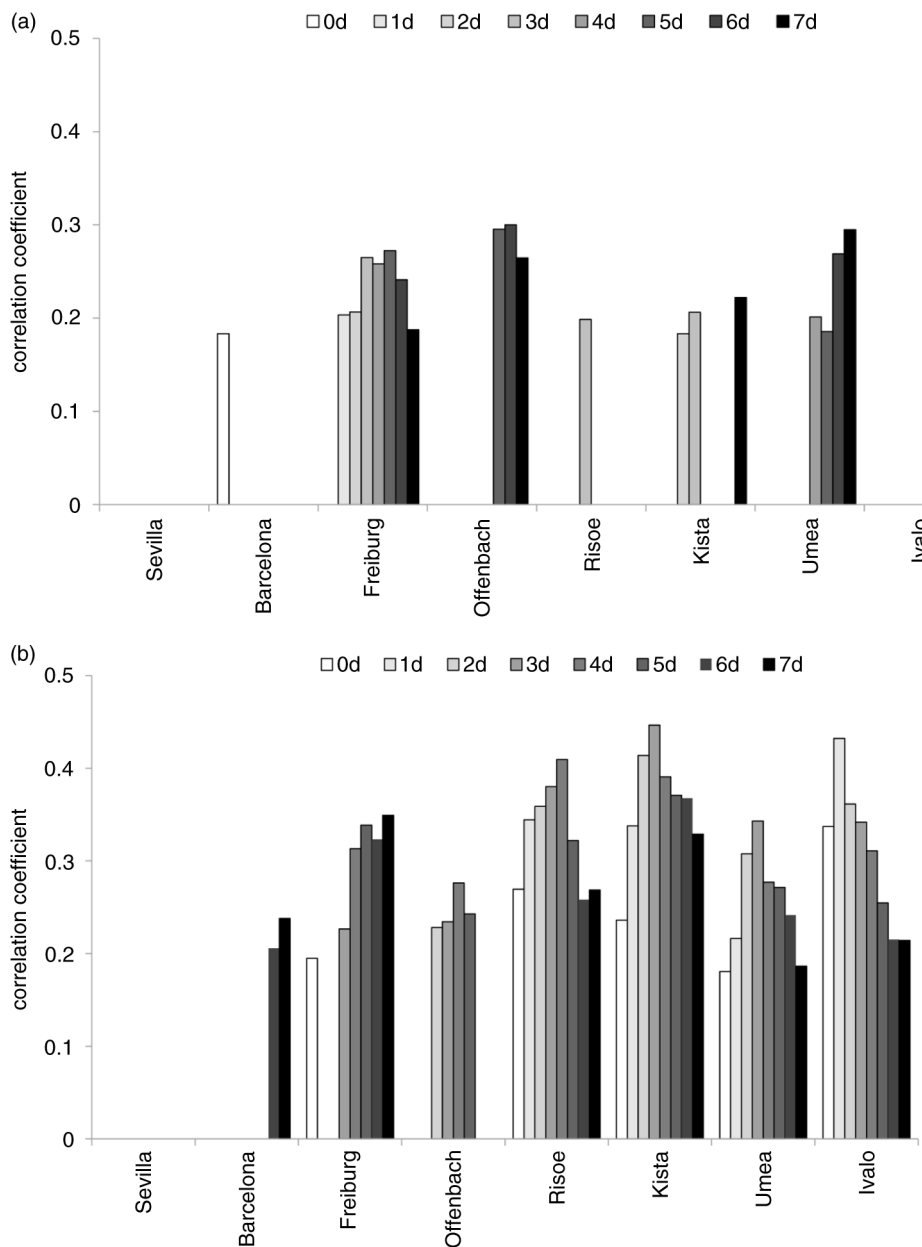


Fig. 8. Winter (a) and summer (b) significant correlation coefficients between <sup>7</sup>Be and TPH with a varying time lag from zero to seven for the selected stations.

variability, with two distinct groups (south and north) in which seasonal differences were also observed.

These results verify the impact of TPH on the <sup>7</sup>Be surface concentrations and also represent the first attempt to understand the variability of that impact. The large number of measuring locations across Europe offers a study platform of great spatial coverage which could pose a basis for future research towards a general understanding of the relationship between the <sup>7</sup>Be surface concentration and its governing mechanisms.

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